

# A note on the statistical distribution of labeled and unlabeled oxygen atoms from N<sub>2</sub>O decomposing on Fe-zeolites and the equilibrium constant between dioxygen isotopomers

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Pirngruber and Roy [1,2] claim that the statistical distribution of <sup>18</sup>O atoms during the decomposition of nitrous oxide over Fe-MFI labeled by <sup>18</sup>O leads to the release of dioxygen species in equilibrium (or, in some transient experiments, out of equilibrium). The equilibrium constant  $Q = [^{16}\text{O}^{18}\text{O}]^2 / ([^{16}\text{O}_2][^{18}\text{O}_2])$  is equal to 4 at temperatures common in catalyzed decomposition of nitrous oxide. In paper [2] both denominator terms are squared, probably by a printing error. The (non) equilibration state has led Pirngruber and Roy to speculations on the reaction mechanism. However, we would like to point out that the recombination of statistically distributed <sup>18</sup>O and <sup>16</sup>O does not lead to the desorption of dioxygen isotopomers in equilibrium. Let us consider that there are, e.g., 3 atoms of <sup>18</sup>O and 7 atoms of <sup>16</sup>O; the probability of the release of <sup>18</sup>O<sub>2</sub> is 0.06, of <sup>16</sup>O<sup>18</sup>O 0.47 and the same for the release of <sup>16</sup>O<sub>2</sub>. The Q value is clearly far off 4, similarly as for different <sup>16</sup>O/<sup>18</sup>O ratios. The interaction of dioxygen species with active oxygen atoms in the zeolite has to be clearly involved. The authors [1,2] have not found such a reaction, most probably because of inappropriate dioxygen and <sup>18</sup>O ratios. Contrary to this, we have reported on a rapid equilibration between oxygen from nitrous oxide labeled

by <sup>18</sup>O and unlabeled dioxygen over Fe-ferrierites below 300 °C [3]. As the isotopic exchange of only <sup>18</sup>O<sub>2</sub> with Fe-ferrierite does not proceed below at least 500 °C, dioxygen molecules in the mixture with labeled N<sub>2</sub>O must react with active oxygen atoms formed from nitrous oxide. Similarly, the equilibrium in isotopomers of dioxygens released from decomposing N<sub>2</sub><sup>18</sup>O (without additional dioxygen in the gas phase, which does not affect the reaction rate) is also rapidly established. We have therefore suggested that this equilibration proceeds between primarily released <sup>18</sup>O<sub>2</sub> and oxygen of the zeolite by a single mechanism – the isotope exchange of 1 atom from the dioxygen molecule for one atom in the surface in one reaction step – for details see [4].

## References

- [1] G.D. Pirngruber and P.K. Roy, *Catal. Lett.* 93 (2004) 73.
- [2] G.D. Pirngruber and P.K. Roy, *Catal. Today* 110 (2005) 199.
- [3] J. Nováková, M. Schwarze and Z. Sobalík, *Catal. Lett.* 104 (2005) 157.
- [4] K. Klier, J. Nováková and P. Jirů, *J. Catal.* 2 (1963) 479.

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